

# Patterns and impacts of ocean warming and heat uptake

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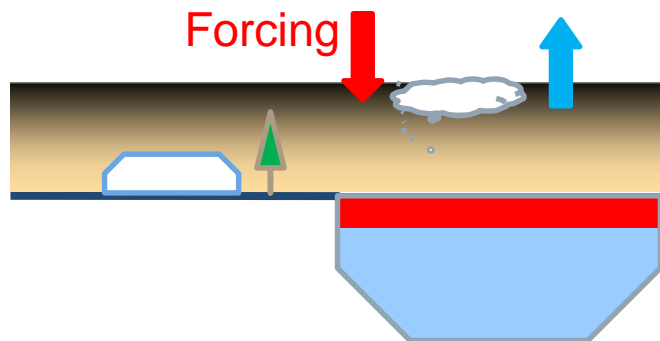
- Ocean warming & circulation change
- Ocean heat uptake & meridional overturning circulation
- Global warming hiatus



# Planetary Energy balance

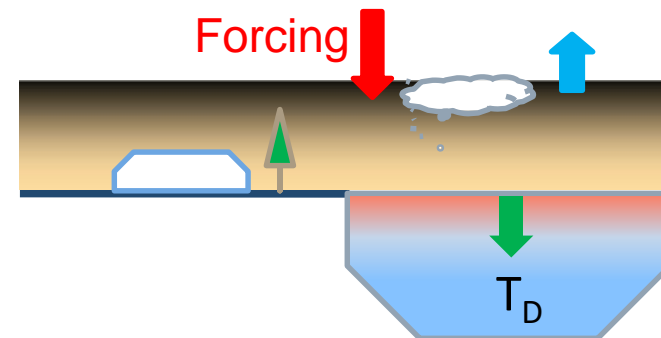
$$H_t = NTO_A = F - \lambda T$$

Ocean heat content change  
TOA radiative imbalance  
Radiative forcing  
Climate feedback



No exchange w/ deep ocean  
→ TOA energy flux = 0 in ~10 yrs  
→ **Equilibrium response**

*Ocean heat uptake slows sfc warming*

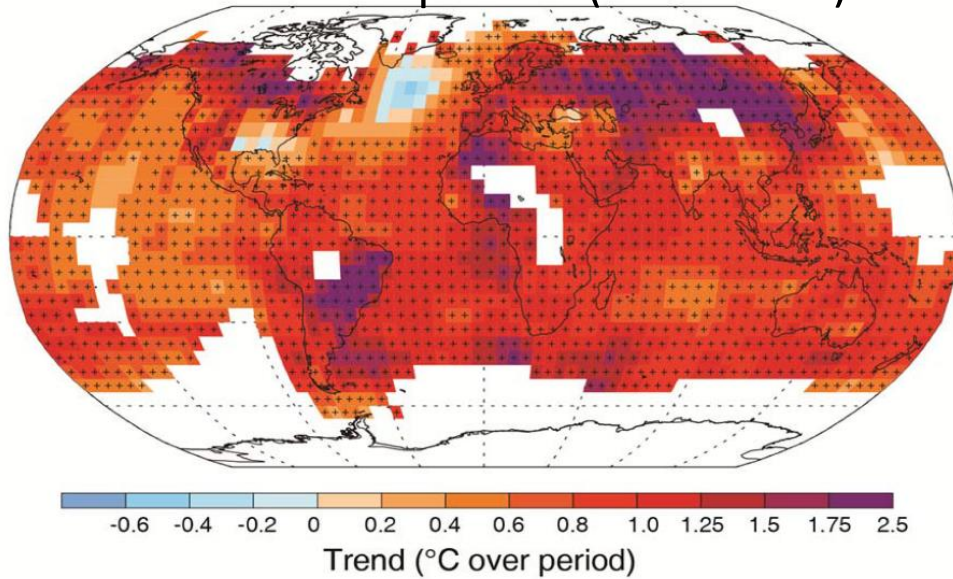


Exchange w/ deep ocean  
→ Ocean heat uptake = TOA energy flux  
→ Smaller **transient response**

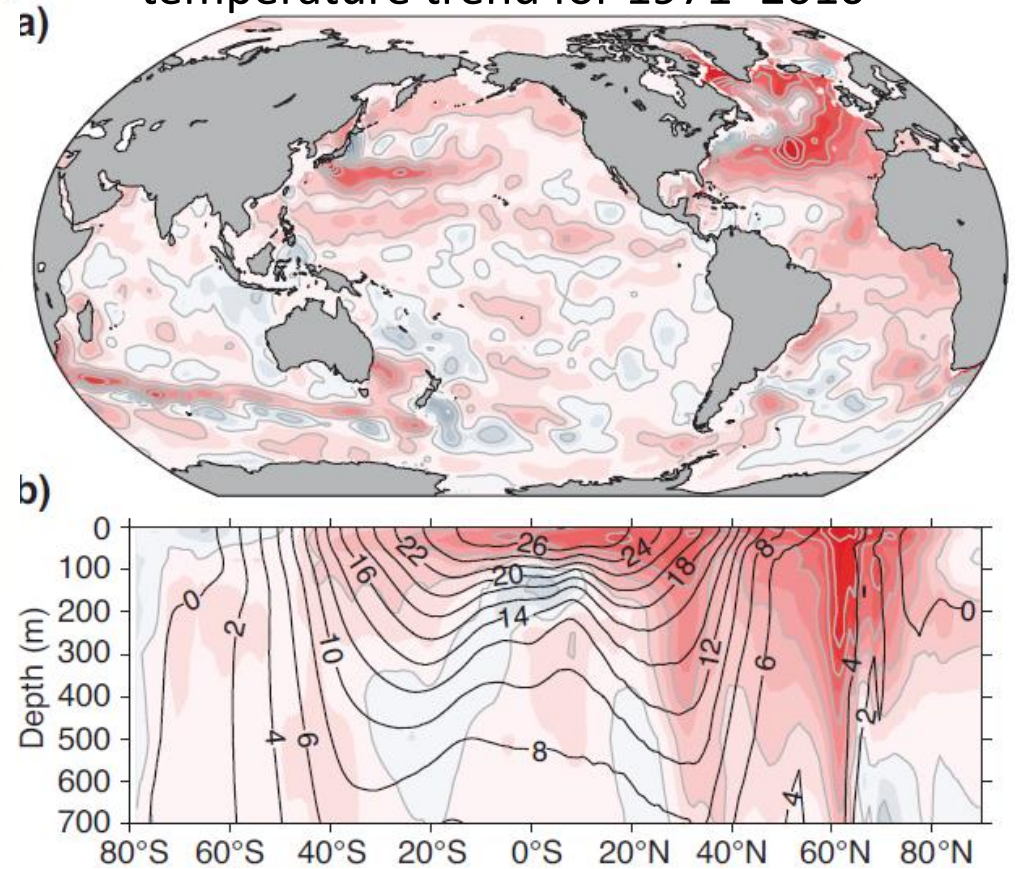
# Historical warming

Ocean warming stores energy and slows atmospheric warming by ~50%.

Sea sfc temp trend (1901-2012)



0–700 m depth-averaged temperature trend for 1971–2010



Most ocean is stably stratified  
→ warming is most confined  
above the thermocline.

Zonal-mean ocean temp change & mean

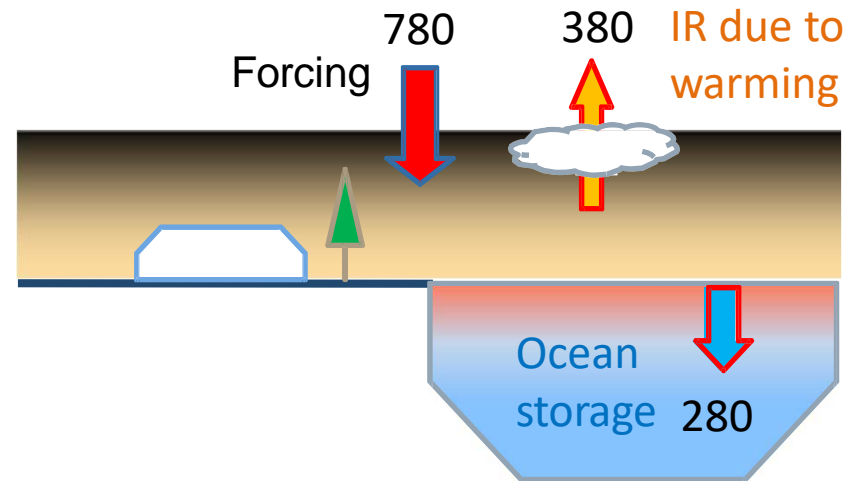
## Planetary energy balance

- Ocean stores 93% of Earth's energy gain.
- Anthropogenic radiative forcing is needed to balance increased radiation into space and ocean heat uptake.

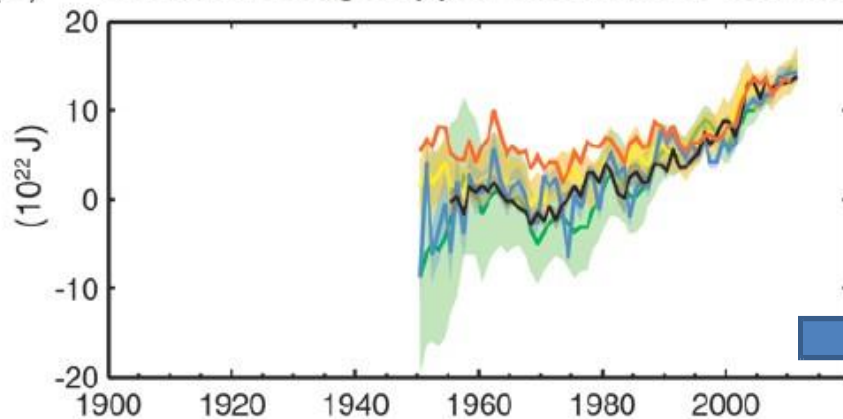
### Ocean heat uptake

+ ) increased outgoing IR

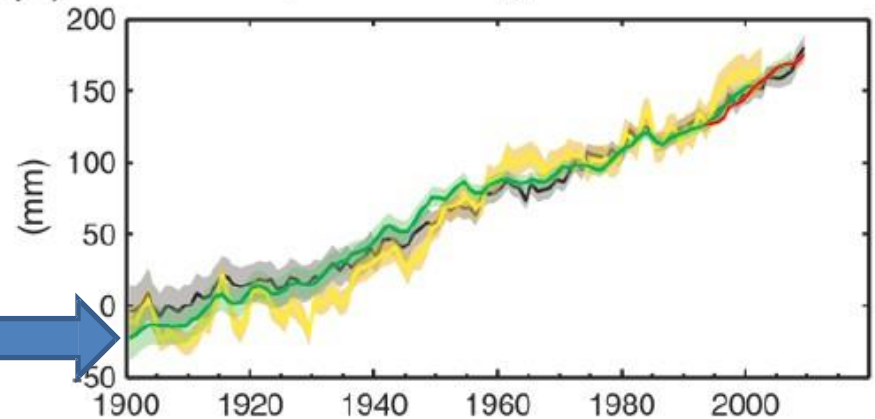
= radiative forcing (CO<sub>2</sub> ...)



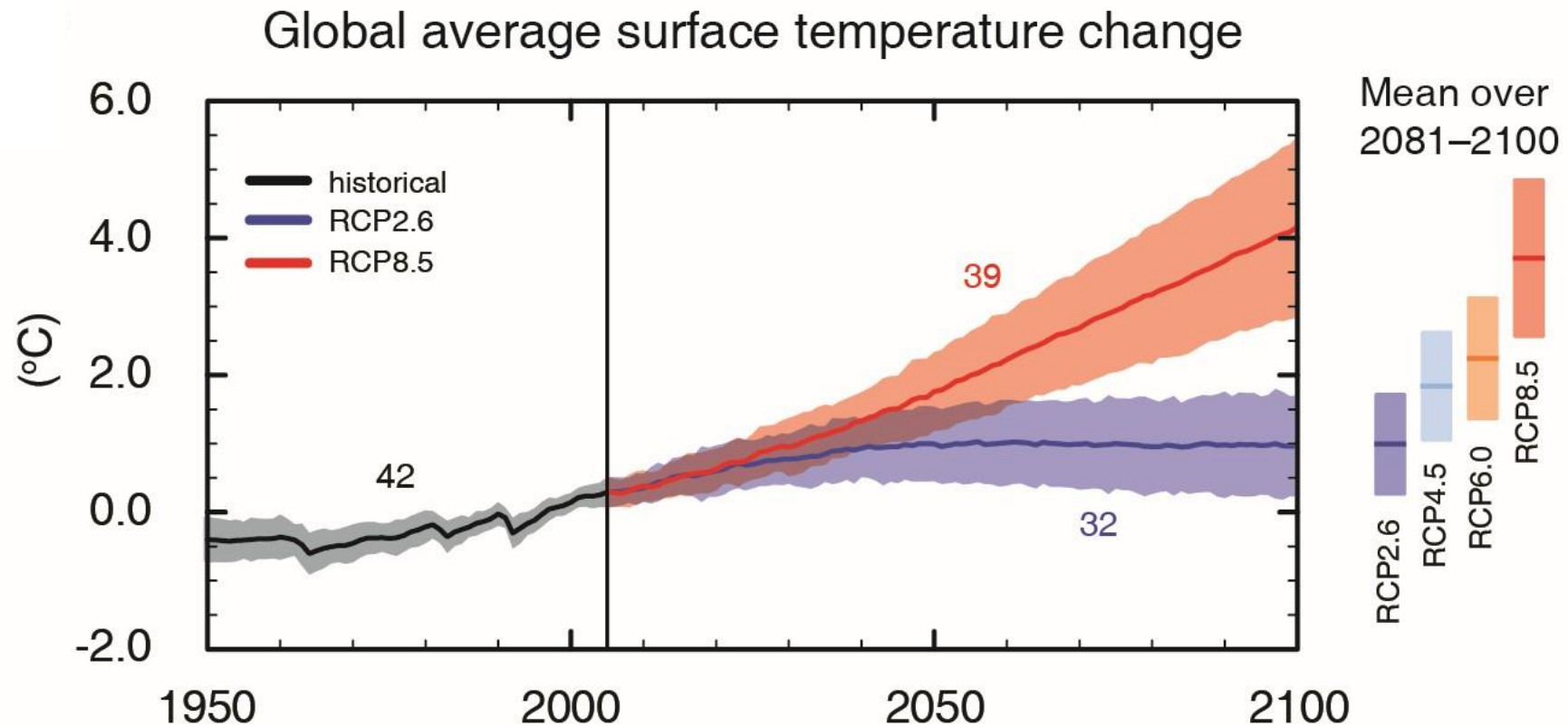
(c) Global average upper ocean heat content



(d) Global average sea level



# Climate Projections



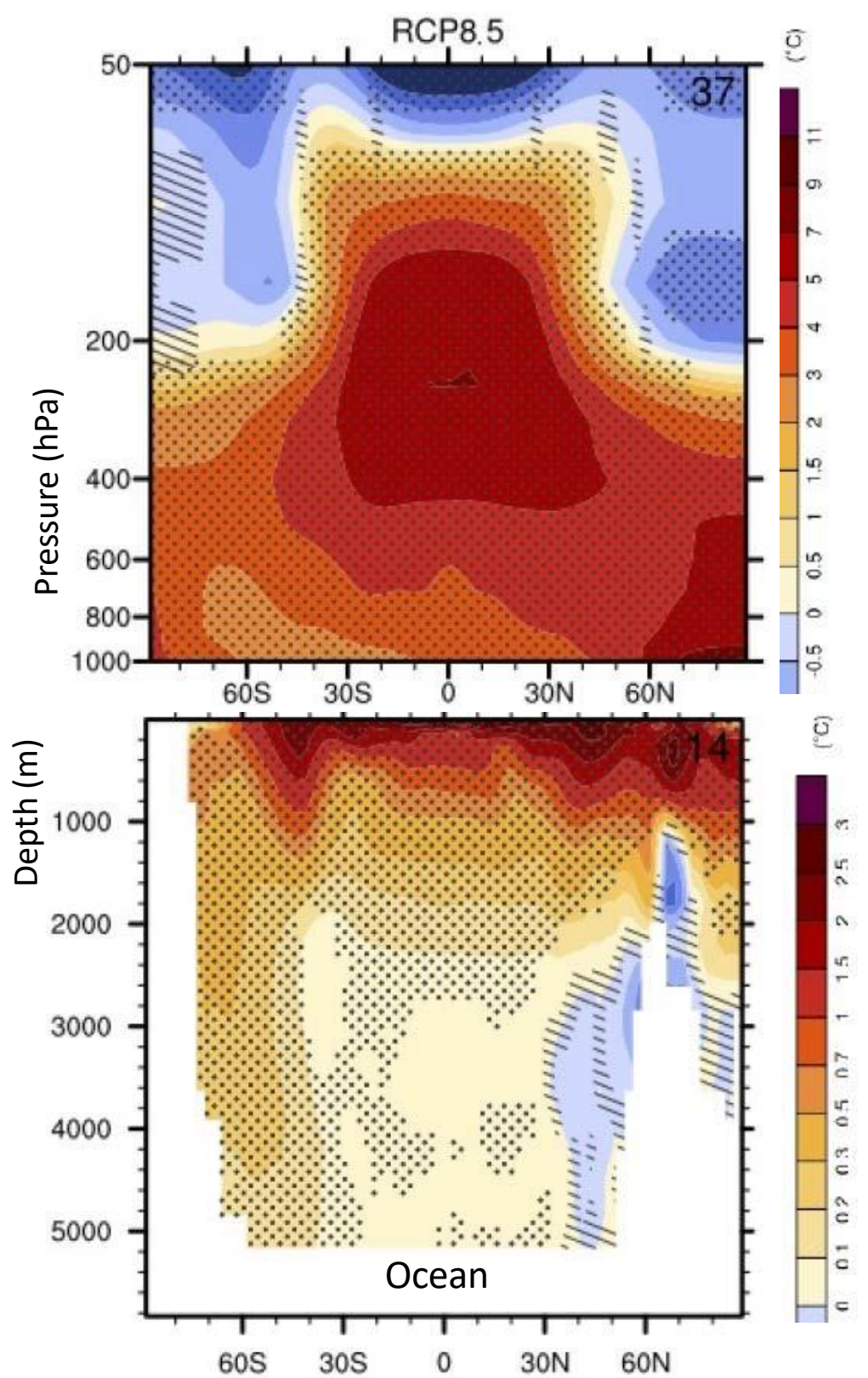
Global surface temperature change for the end of the 21st century is *likely* to exceed 1.5°C relative to 1850.

We are in control of the Earth's thermostat.

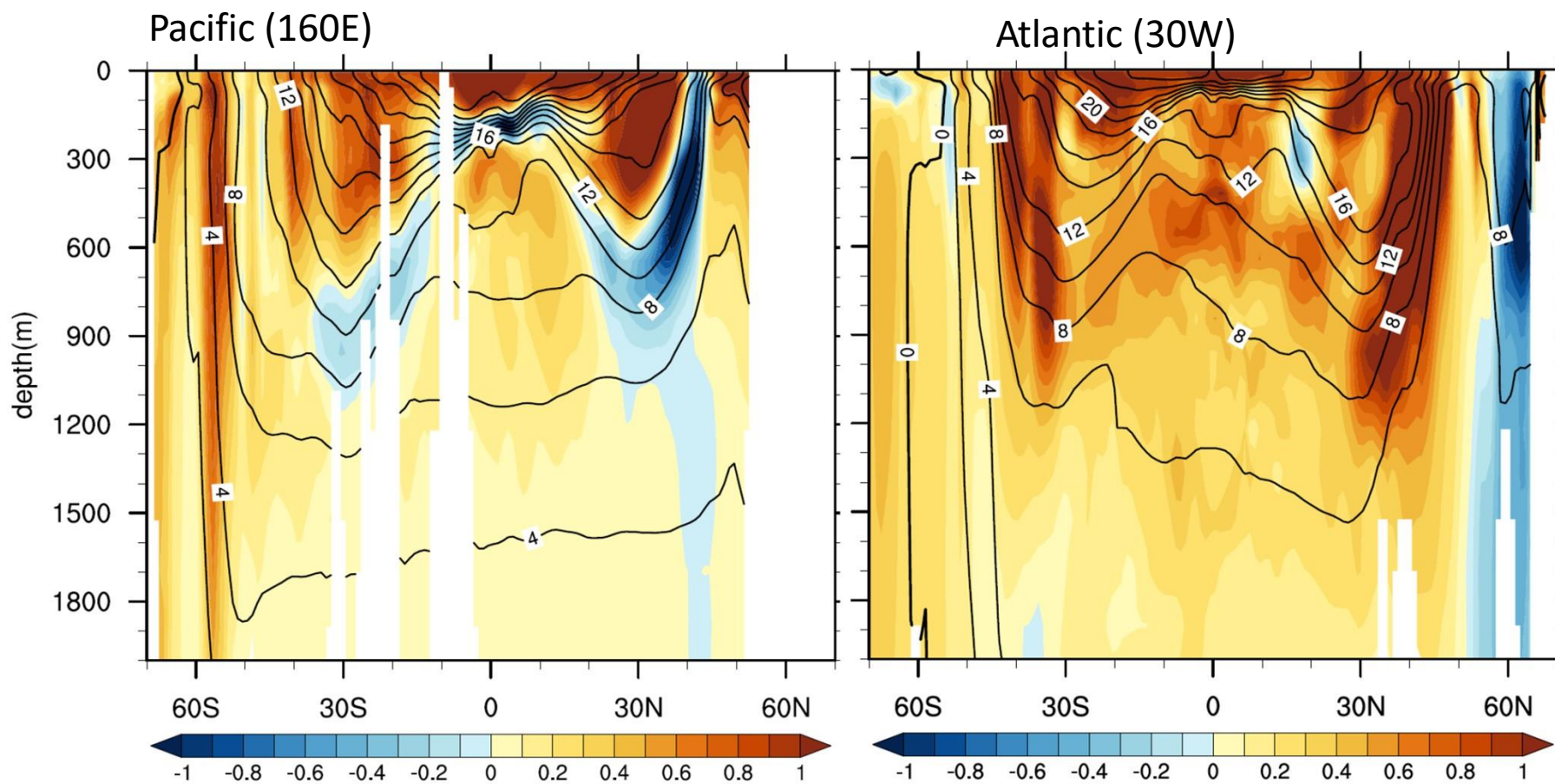
# Future projections

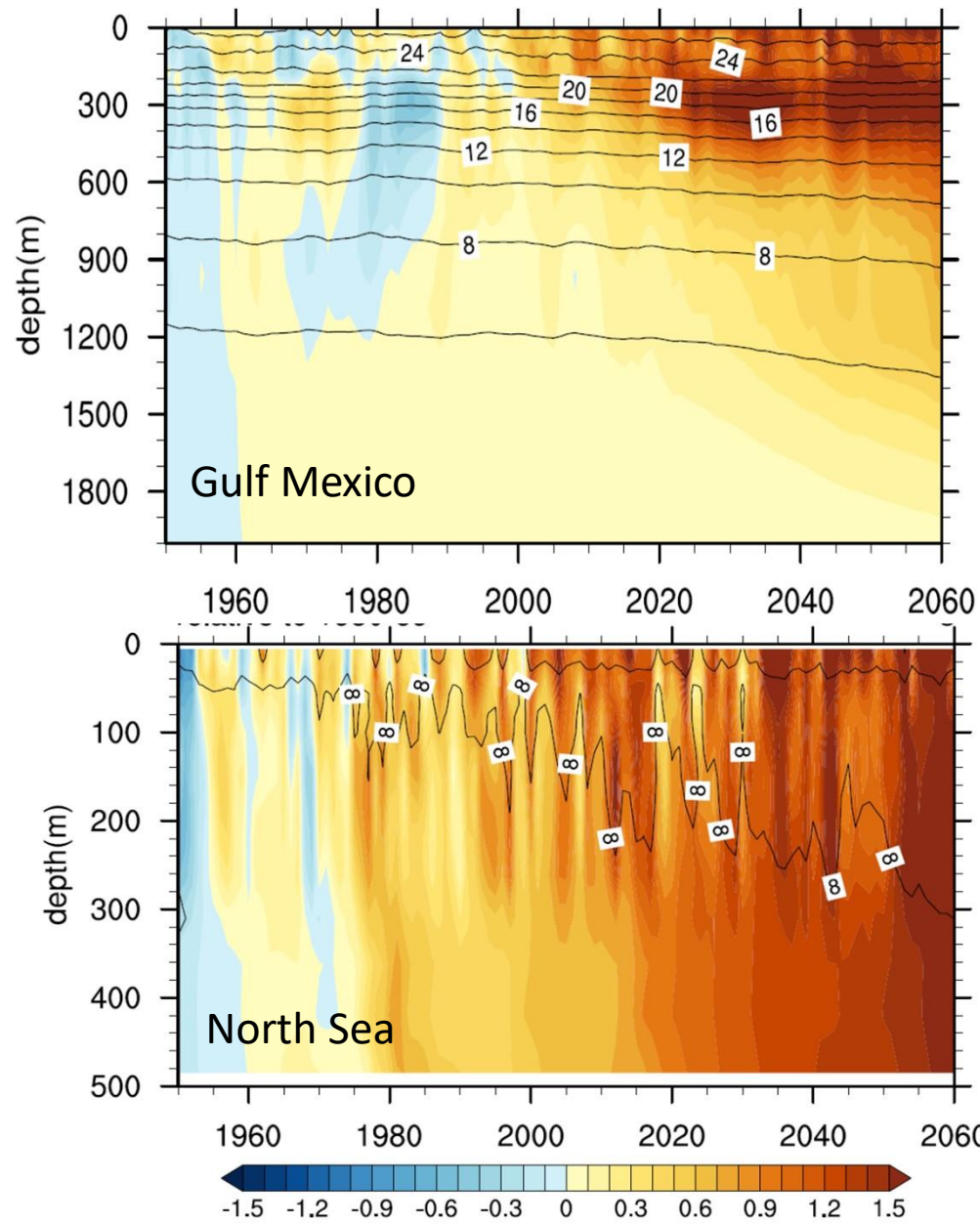
Ann-mean emp change (2081-2100, RCP8.5)

IPCC AR5 (2013)



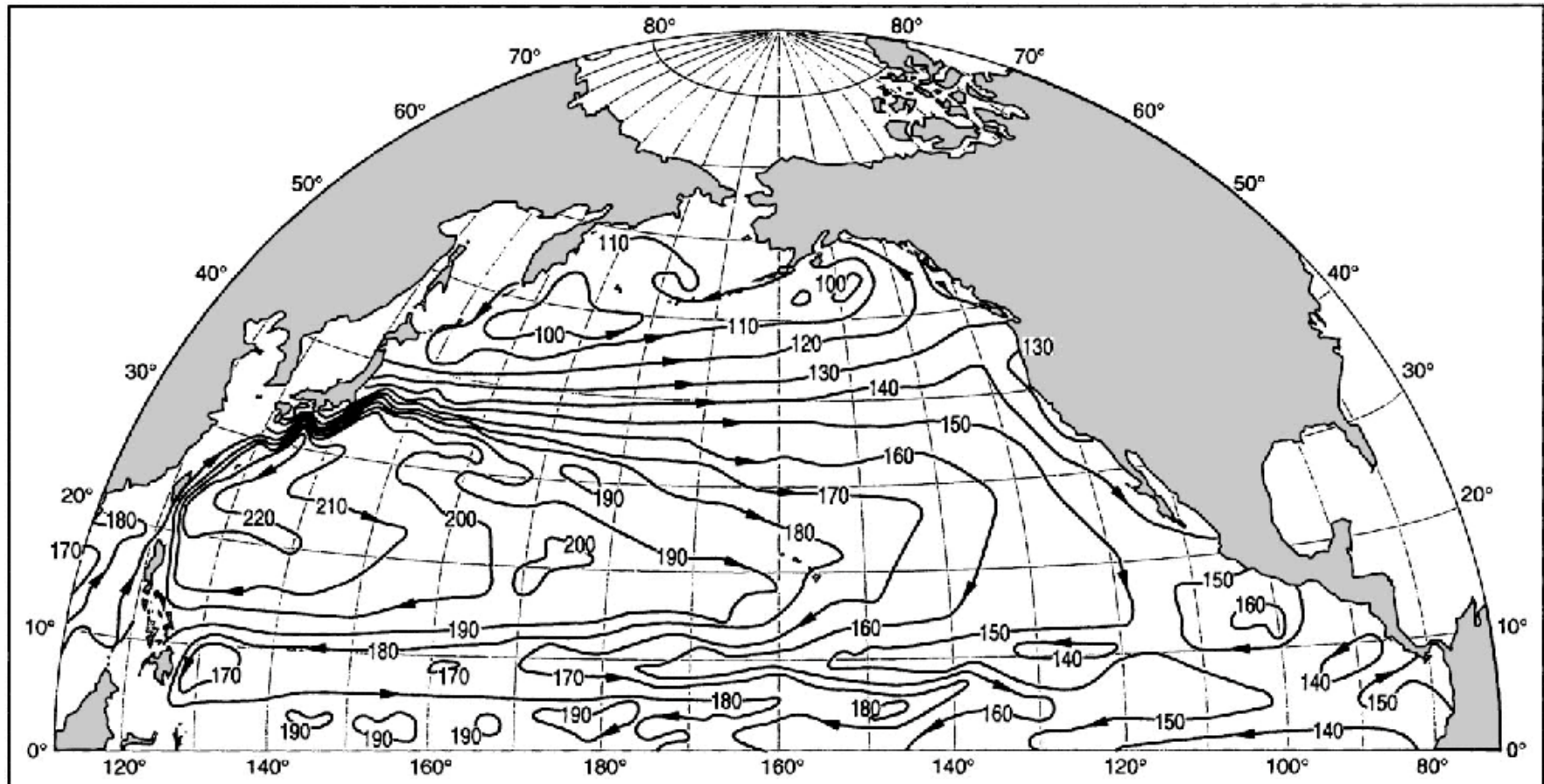
Temp: mean & change from 1985-2005 to 2040-60



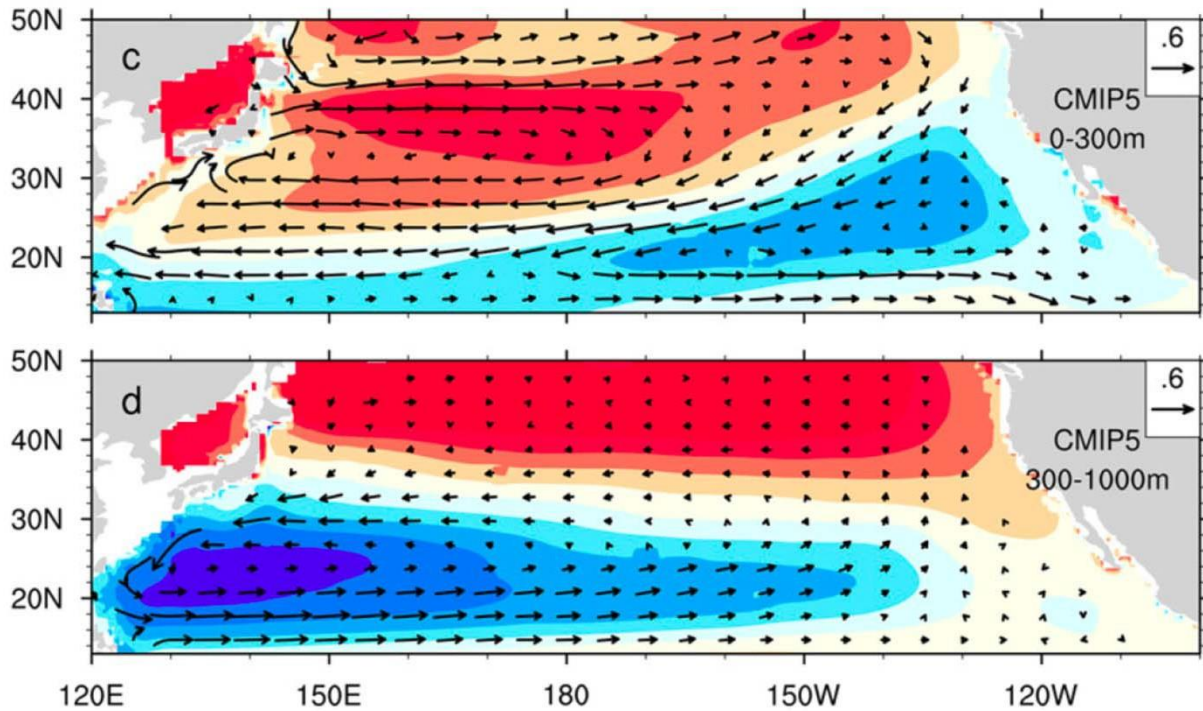


Ocean temp ( $^{\circ}\text{C}$ ): mean & change from 1905-69



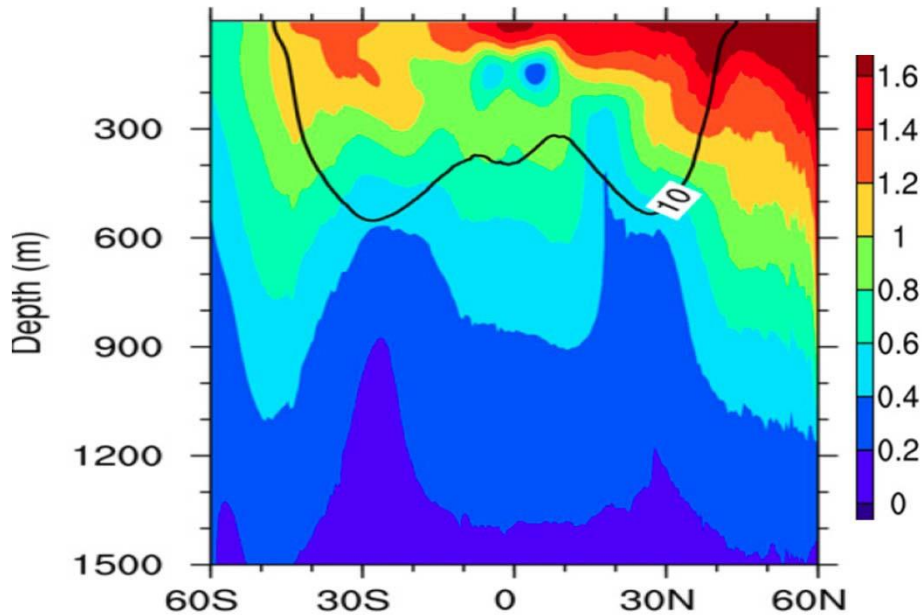


North Pacific surface circulation



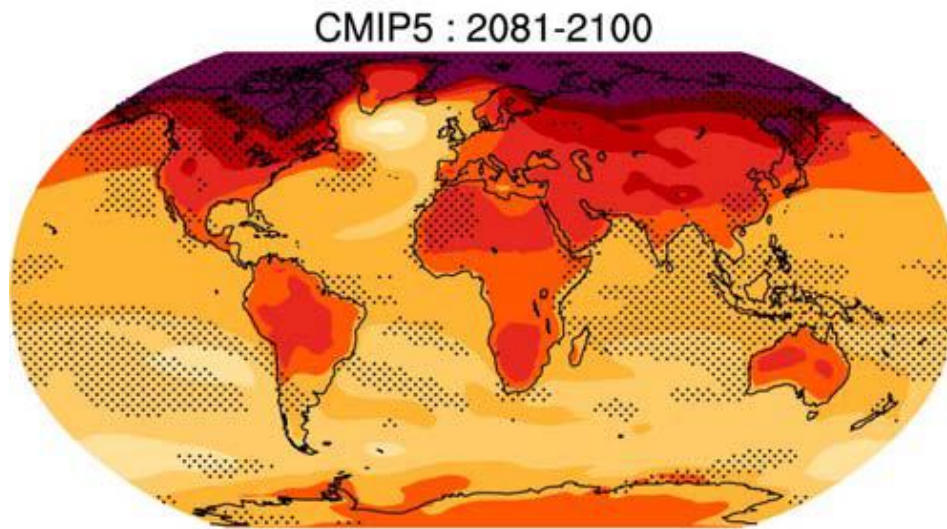
Increased temp stratification  
 ↓  
 Increased upper circulation  
 ↓  
 Weakened lower circulation

Changes in current velocity & streamfunction.

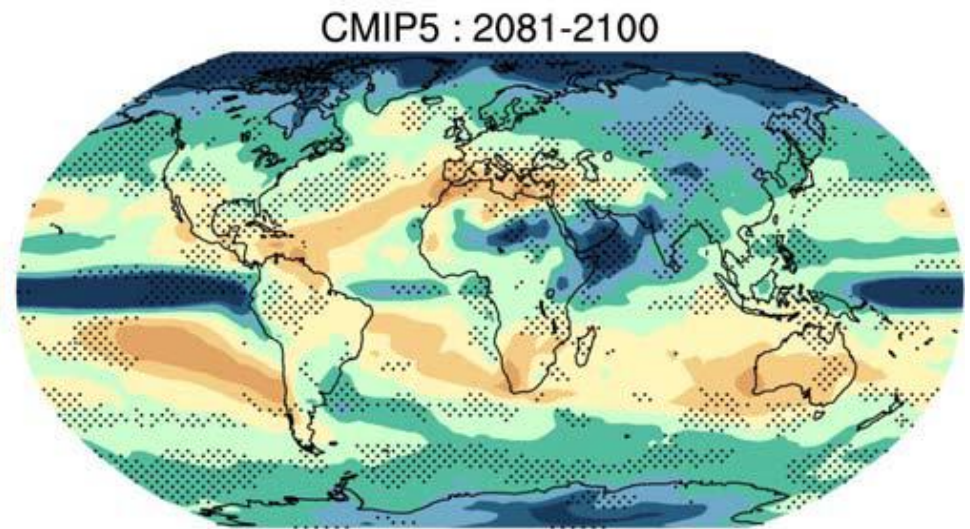
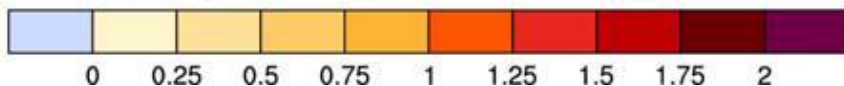


Zonally averaged temperature change ( $^{\circ}\text{C}$ ; 2076-2100 minus 1976-2000) for the Pacific from the CMIP5 ensemble of RCP4.5.

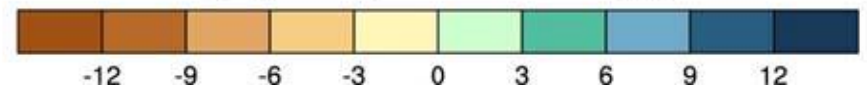
- Temperature rises everywhere.
  - Precipitation increases in some regions and decreases in some others (variable in space).
- What determines regional patterns of precipitation change?



Sfc temp (°C per °C global mean change)

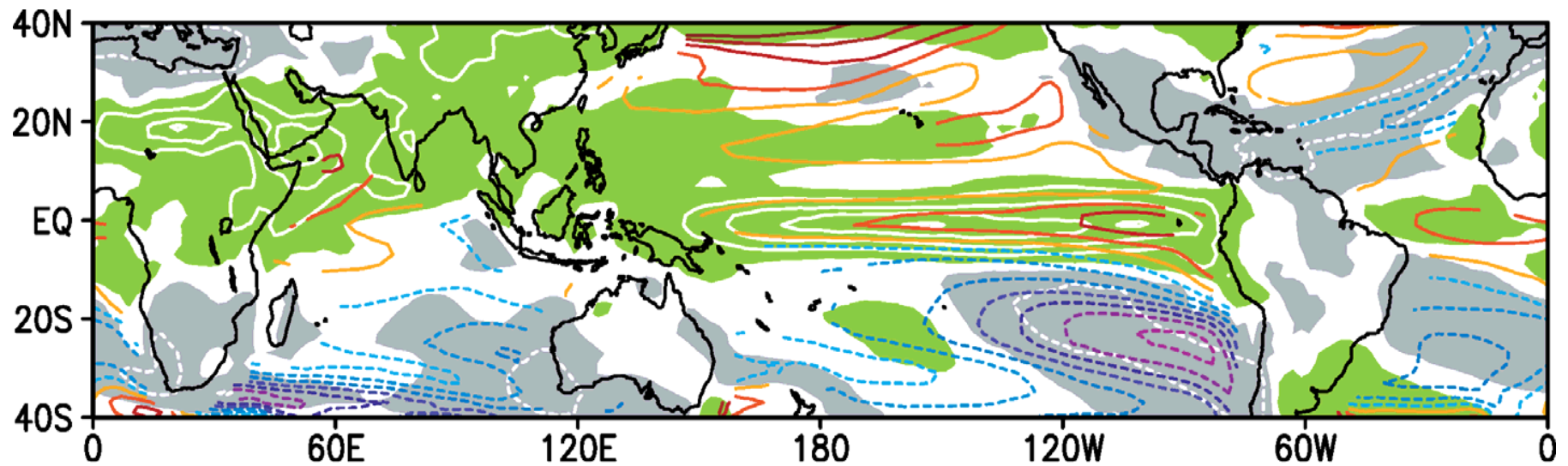


Precip (% per °C global mean change)



## Surface warming effect on tropical rainfall:

Warmer get wetter,  $r(\delta T, \delta P/P)=0.63$



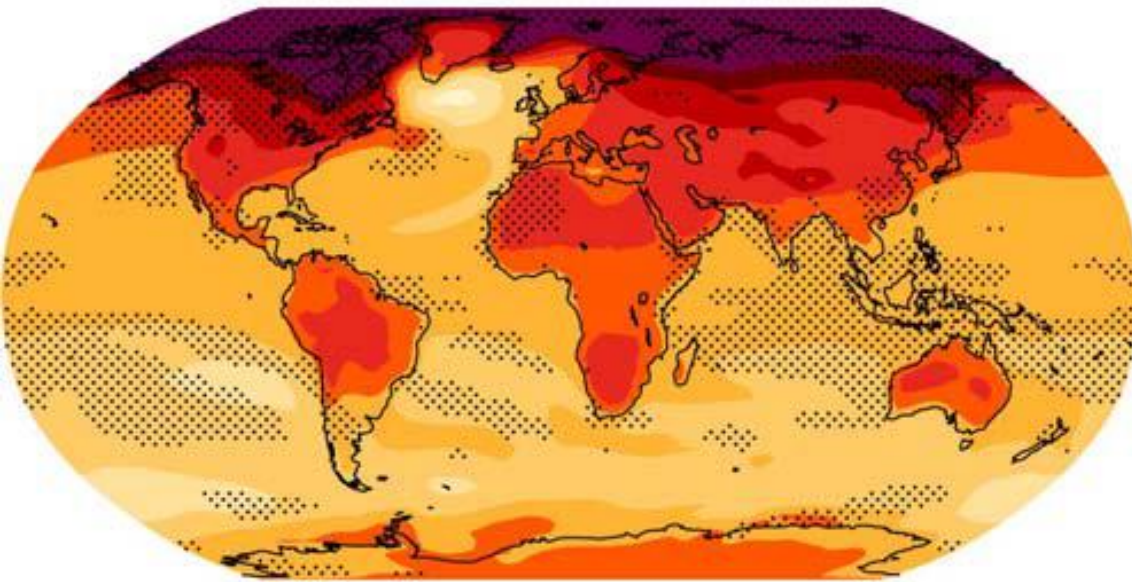
Sea surface warming (color contours) and precipitation change (color shading & white contours) over the 21st century (IPCC AR5, 2013)



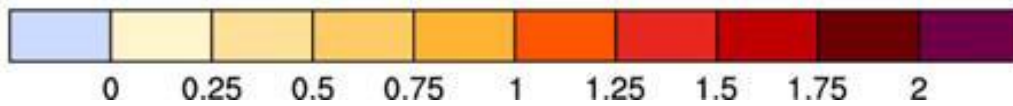
## Summary 1

- Ocean heat uptake and storage is an important component of Earth energy budget.
- Patterns of ocean surface warming affect tropical rainfall, El Nino and storms.
- Rapid ocean warming in and above the thermocline.
- Increased density stratification intensifies the subtropical gyre in the upper layer and slows it down in the lower thermocline.

CMIP5 : 2081-2100



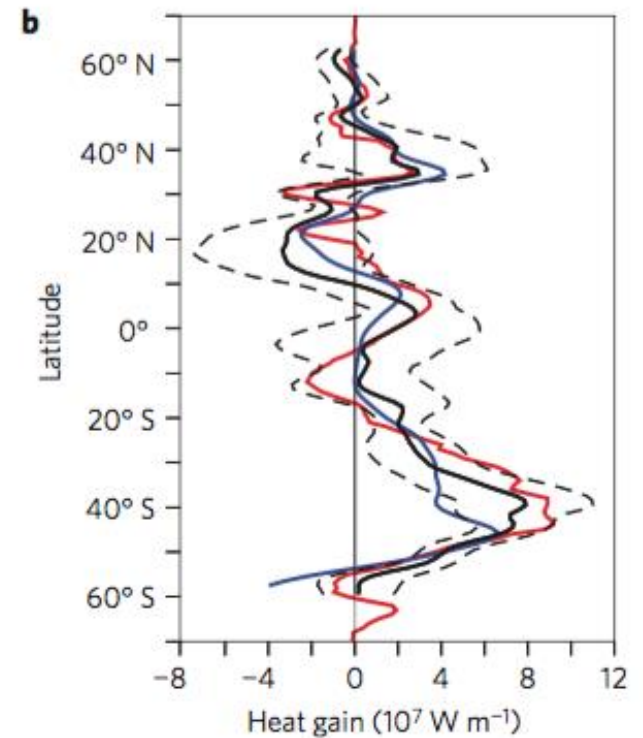
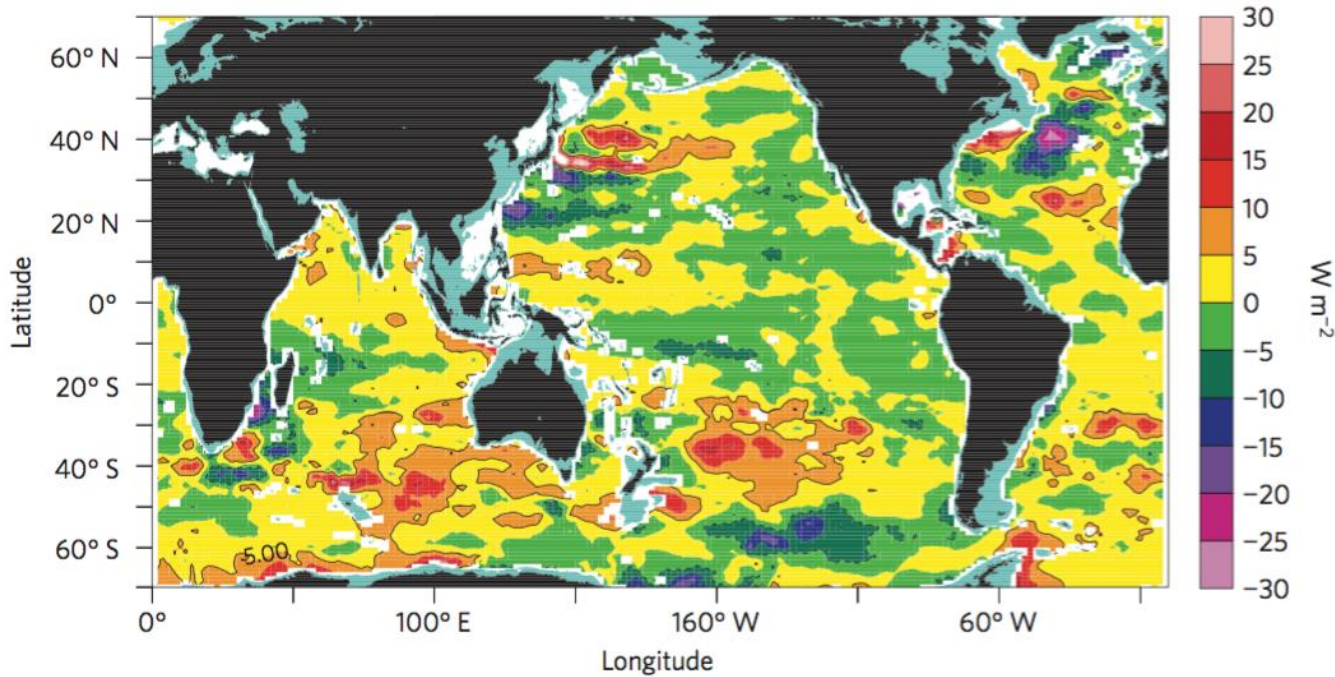
Sfc temp ( $^{\circ}\text{C}$  per  $^{\circ}\text{C}$  global mean change)



- Amplified warming in **Arctic**
- Reduced warming in **Antarctica**

# Uneven ocean heat uptake

## Argo Observations



trend in ocean heat content (0-2000m) for 2006 - 13

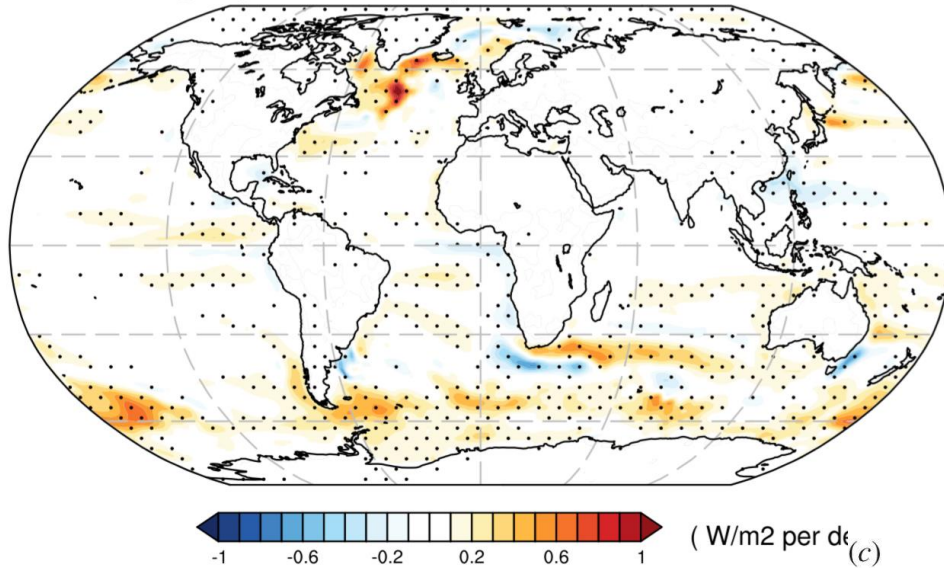
**Most of heat gain (67%~98%) occurred in the S.H. extratropical ocean (south of 20°S)**

(Roemmich et al. 2015, Nat Clim Change)

# Surface heat flux change

CMIP5 ensemble mean

Historical (1861-2005)



# Area-integration over:

NA: 30°N-70°N, 80°W-10°W

SO: south of 30°S

Cumulative Heat Uptake over the 20th century:

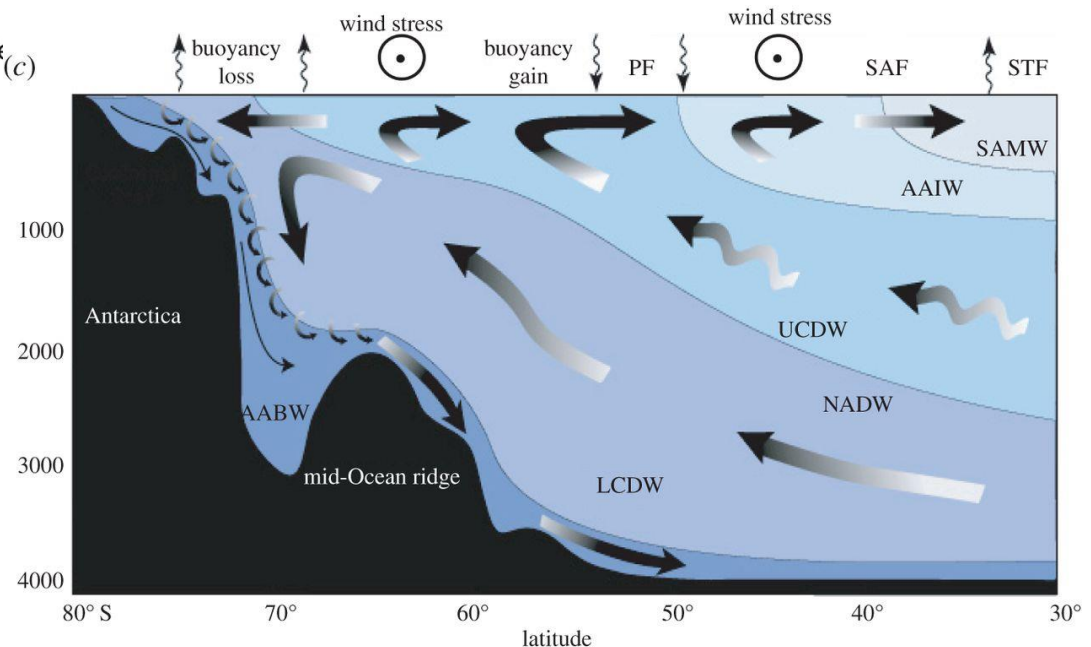
NA: 6%±39%

SO: 72%±28%

## Southern Ocean Heat Uptake

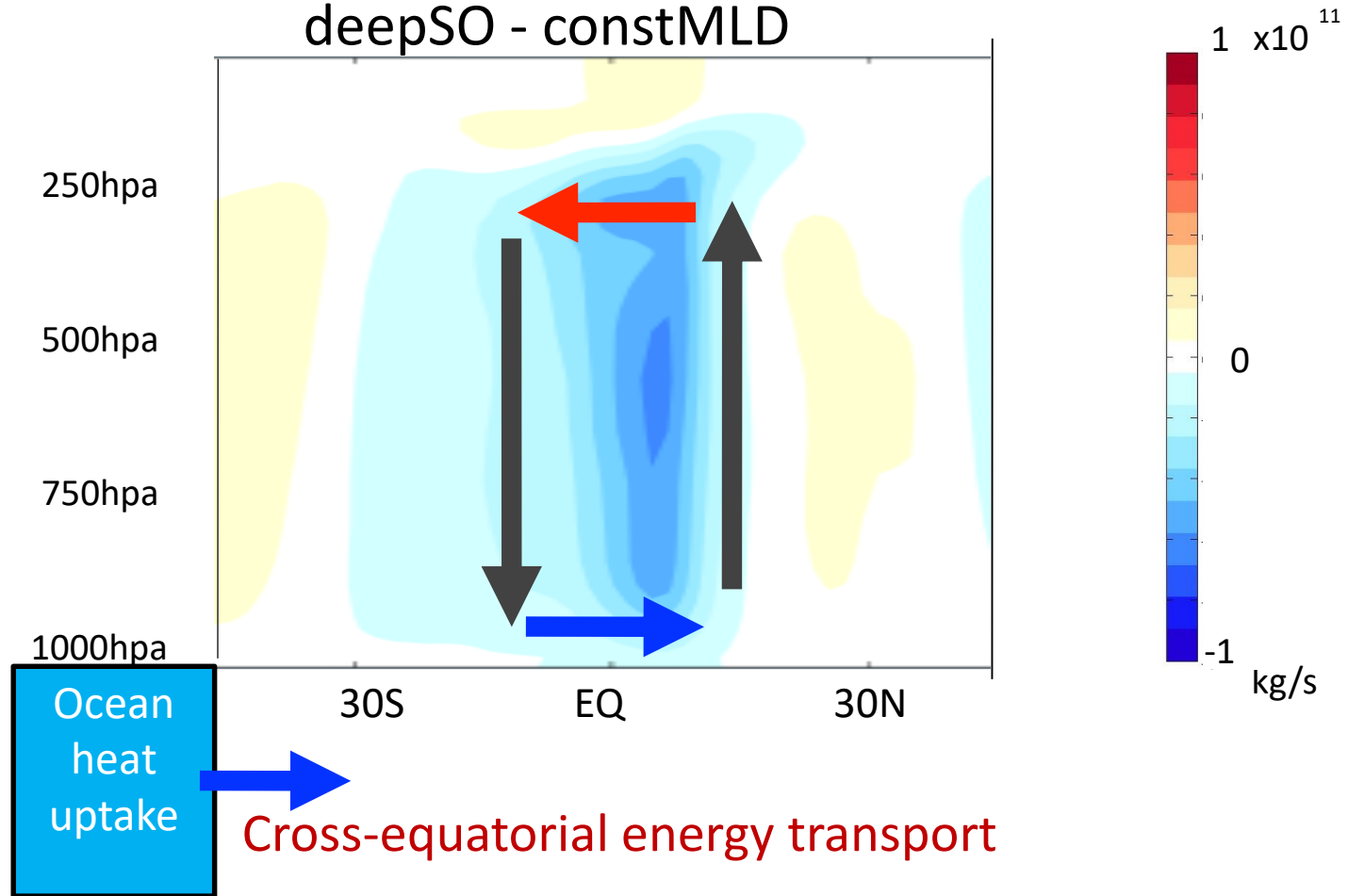
→ Reduced surface warming

(Bryan et al. 1988; Marshall et al. 2014 ...)



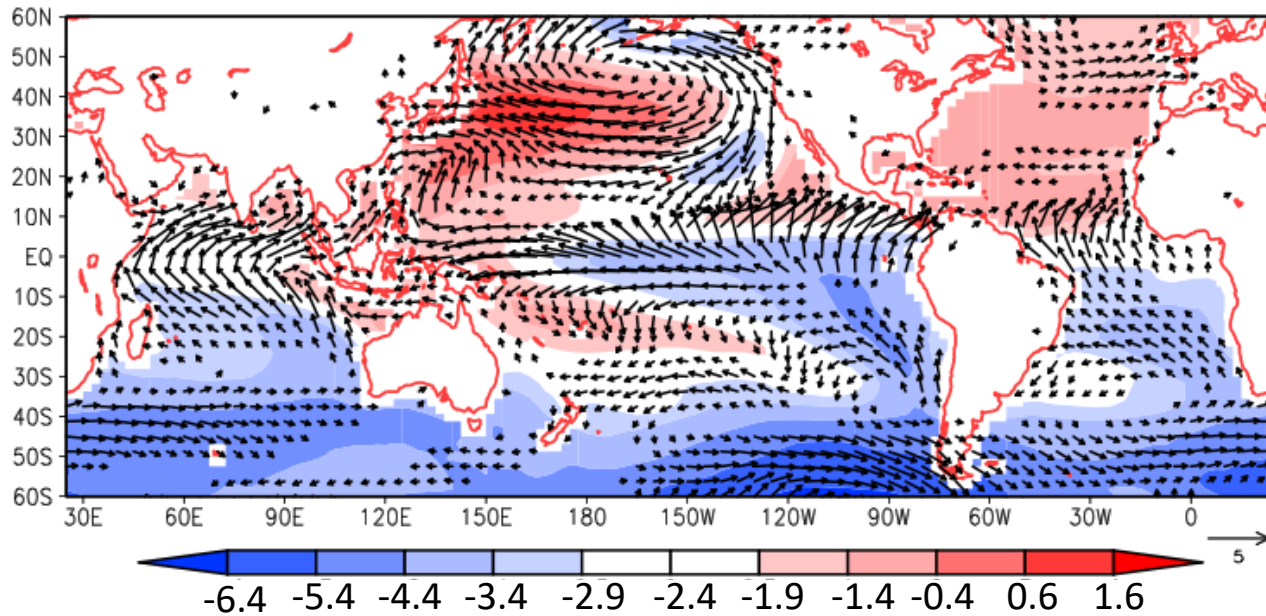


Changes in Streamfunction  
(average of year 10~19 after 4xCO<sub>2</sub>)  
deepSO - constMLD



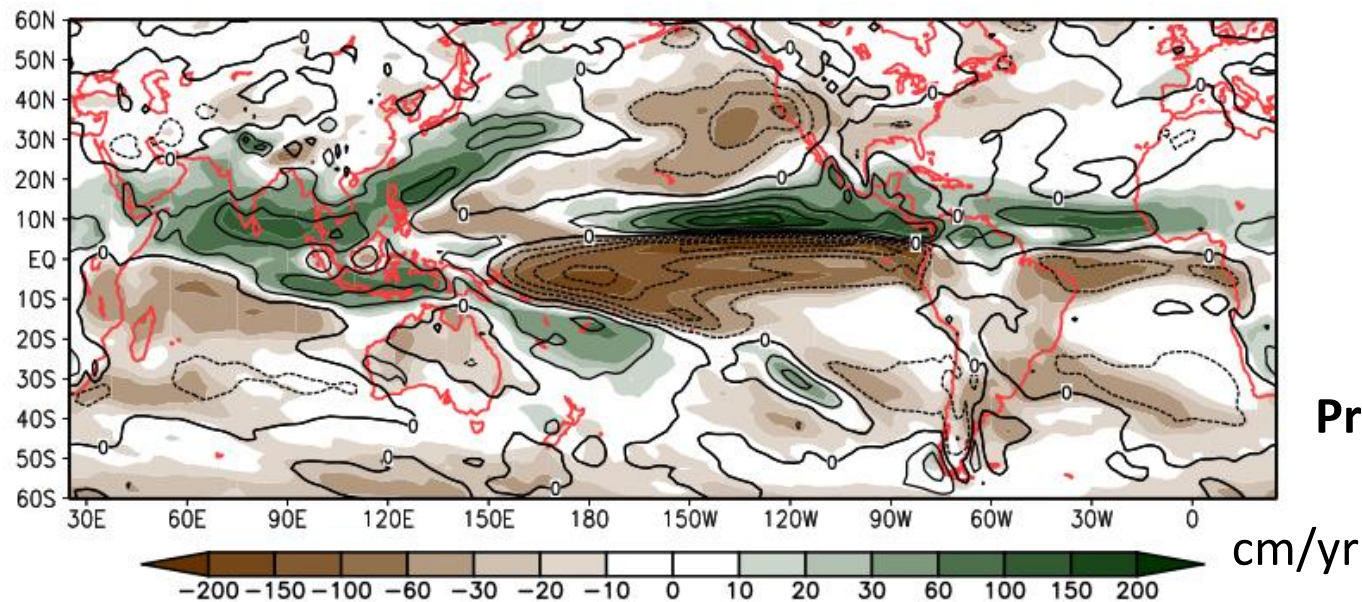
Chiang and Bitz 2005, Kang et al. 2008, 2009; Broccoli et al. 2006, Zhang and Delworth 2005, Frierson and Hwang 2012, Hwang et al. 2013

# deepSO – constMLD Difference (annual mean) (year 10~19 after 4xCO2)



**SST (shading) & Sfc  
Wind (vectors)**

Zonal asymmetry: enhanced cooling in SE Pacific



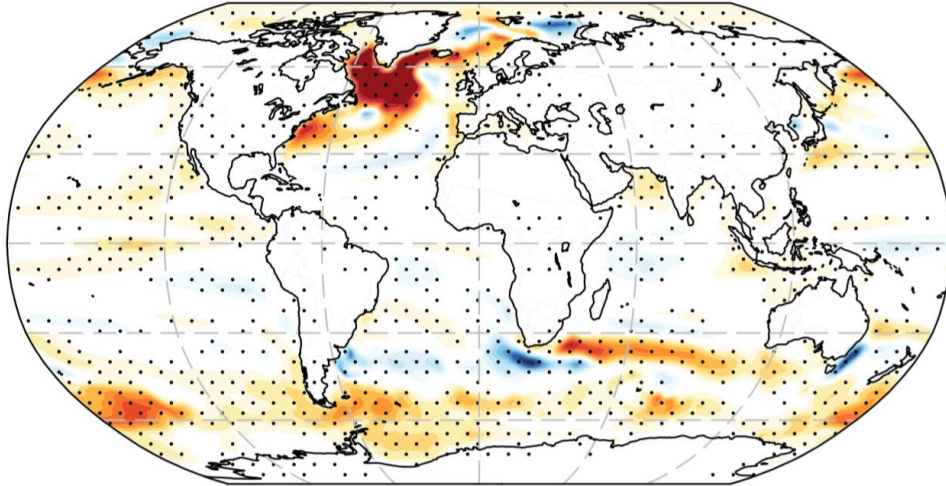
**Precipitation**

cm/yr

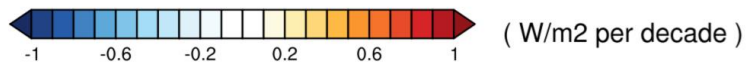
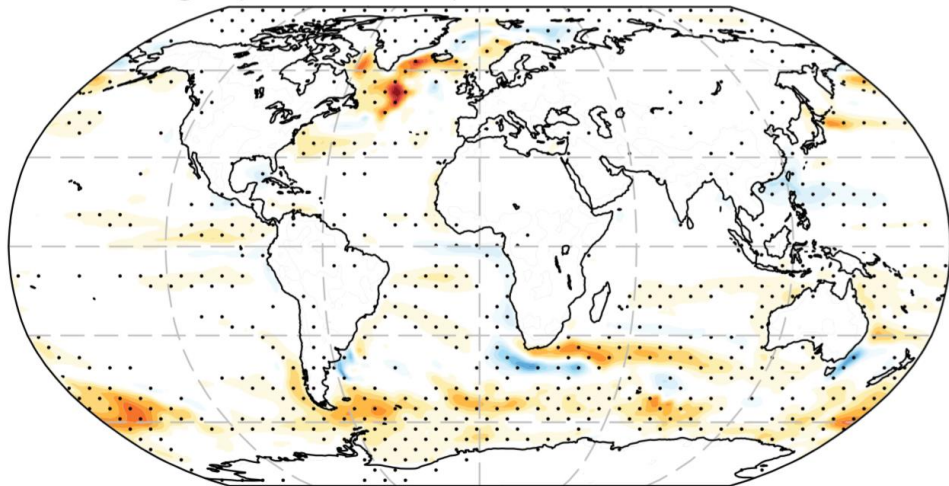
# Response of surface heat flux:

CMIP5 ensemble mean

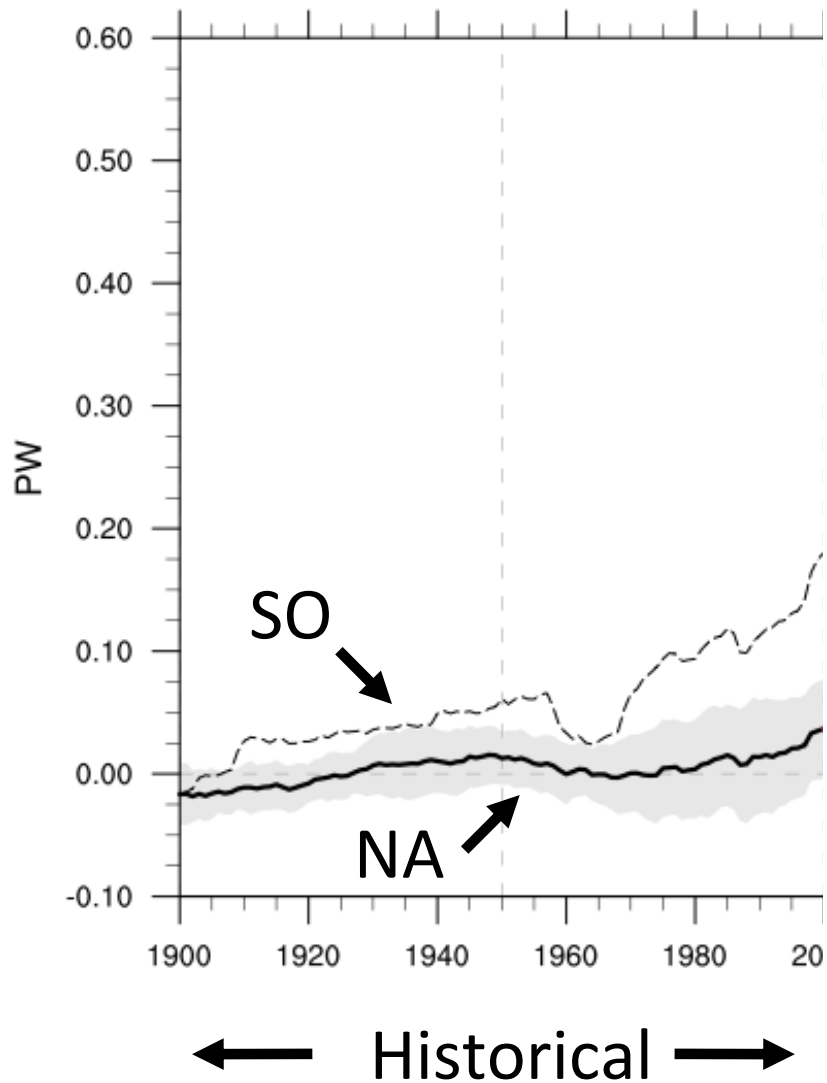
GHG (1861-2005)



Historical (1861-2005)

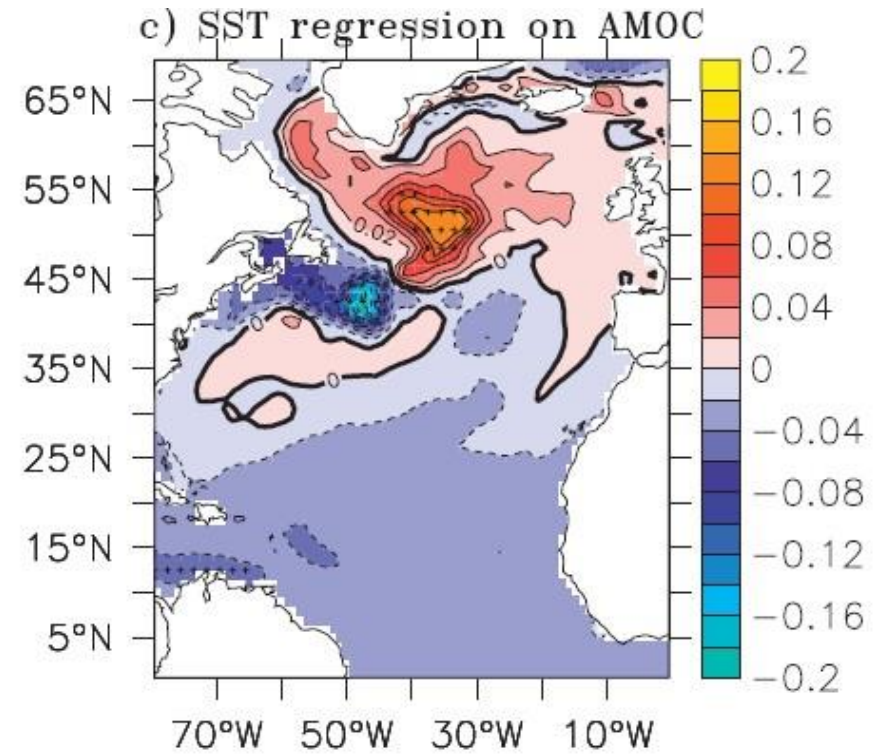
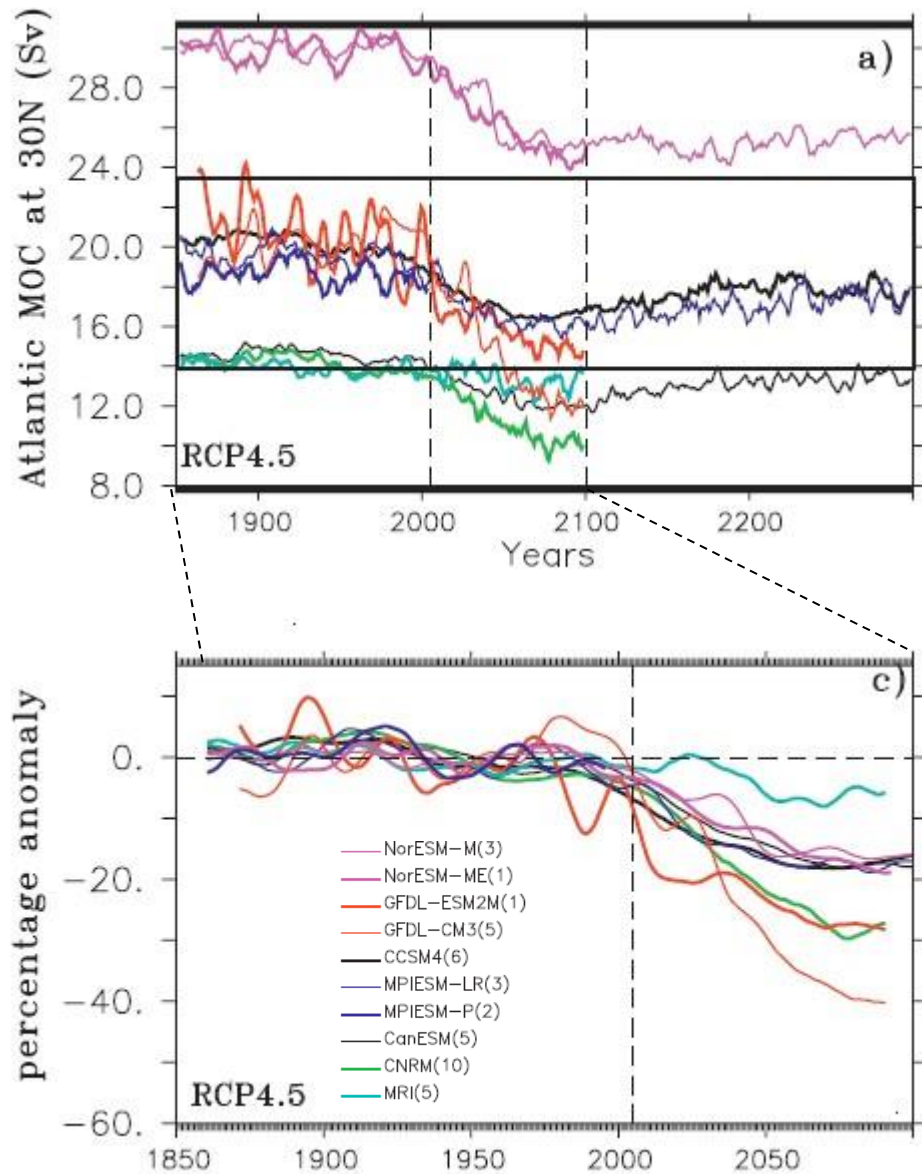


# Area-integrated heat flux



Shi, J., S.-P. Xie, and L.D. Tally, 2018: Evolving relative importance of the Southern Ocean and North Atlantic in anthropogenic ocean heat uptake. *J. Climate*, in press.

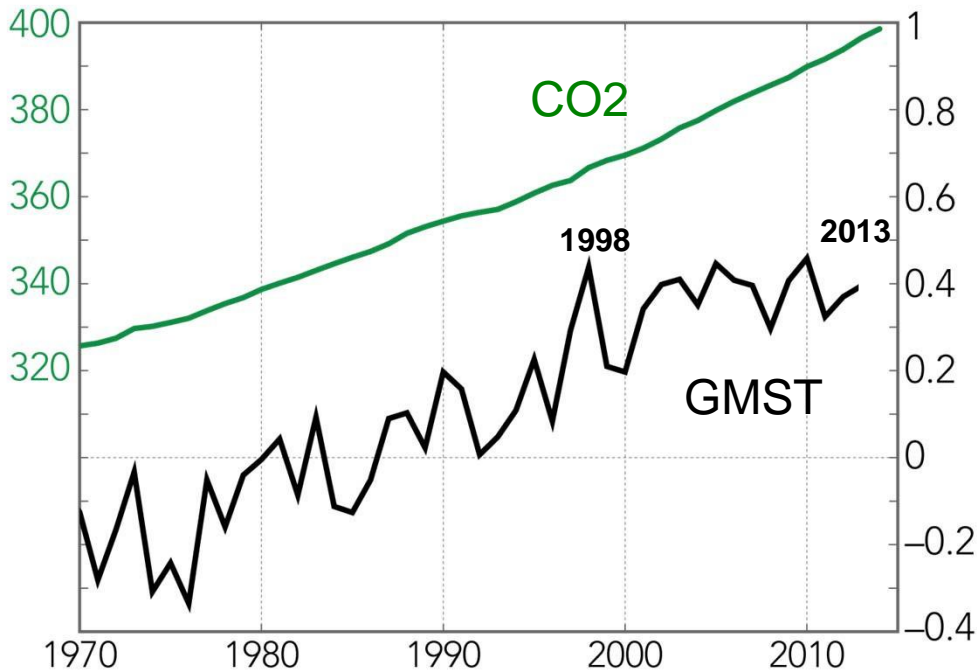
# Atlantic meridional overturning circulation (AMOC) is projected to slow down



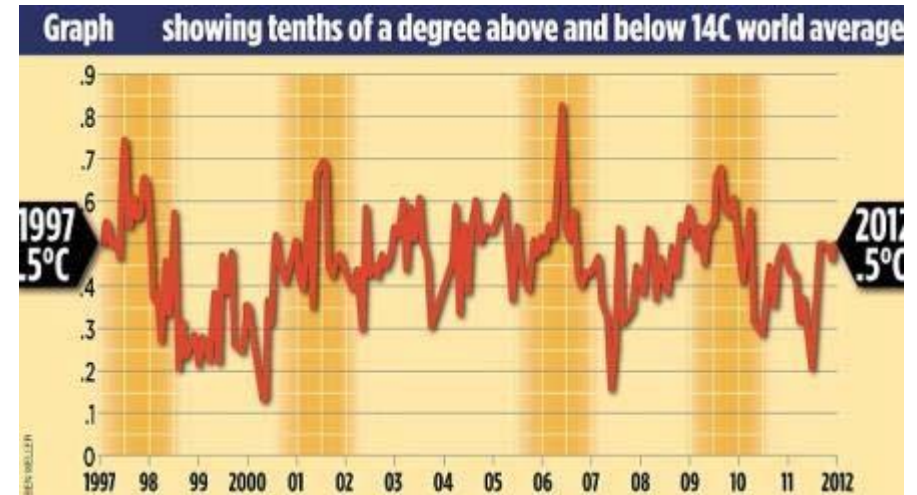
Cheng et al. (2013, J Climate)

## Summary 2

- Historically, most ocean heat uptake took place in Southern Ocean
- Localized heat uptake drives cross-equatorial Hadley circulation response and reorganization of tropical monsoon/rainfall.
- With anticipated decline in aerosols, the North Atlantic contribution to heat uptake is expected to increase,
  - slowing down AMOC
  - Inducing further changes in atmospheric circulation/rainfall.



## Global warming hiatus



*The Mail, UK*

*The New York Times*

**July Set Mark as U.S.'s Hottest Month, August 8, 2012**

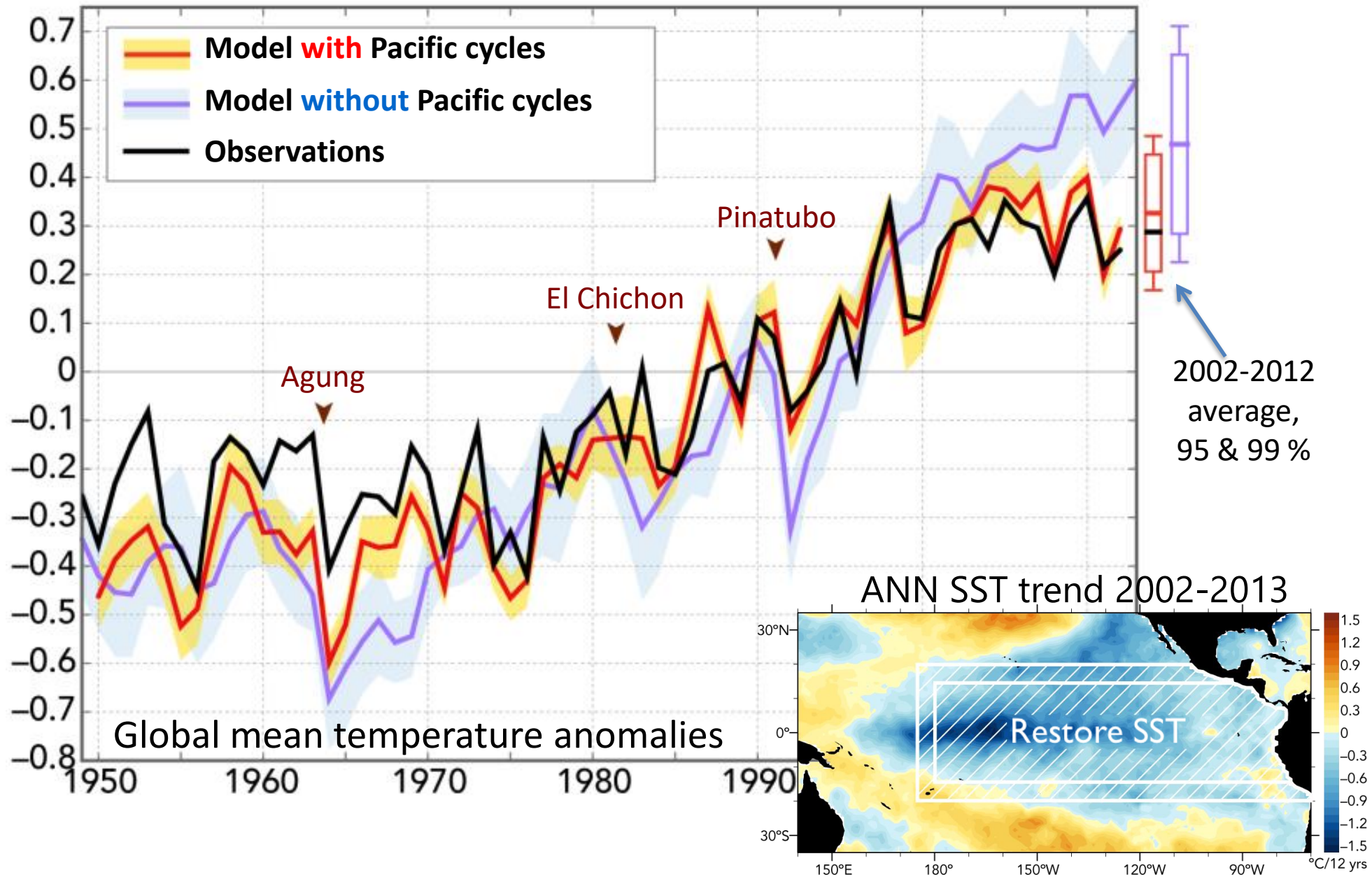
Record heat waves in Europe 2003, Russia 2010, US 2012, China 2013; Record Arctic sea ice loss in 2007 & 2012.

Daily-mean Maua Loa CO2 exceeded 400 ppm for the first time in May 2013.

→ What causes the hiatus?

# POGA (Pacific Ocean-Global Atmosphere) pacemaker run

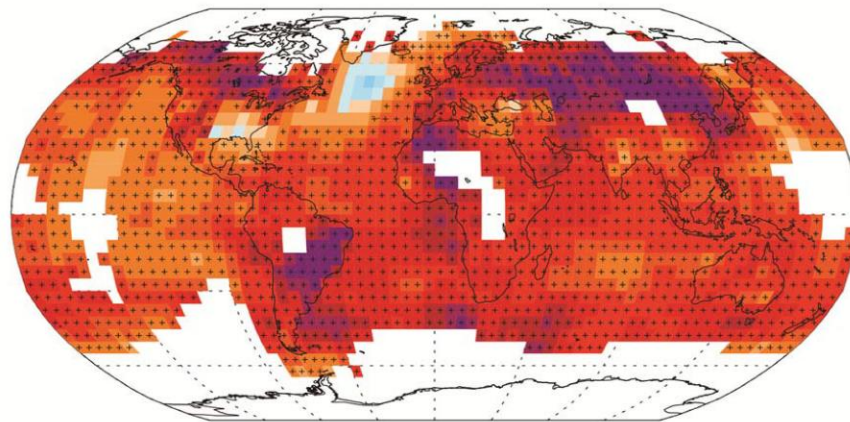
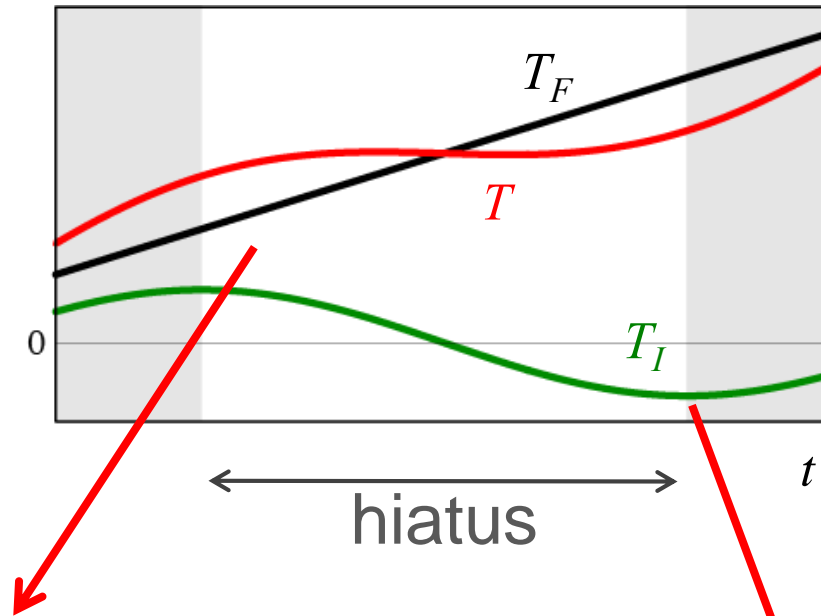
Kosaka & Xie (2013, *Nature*)



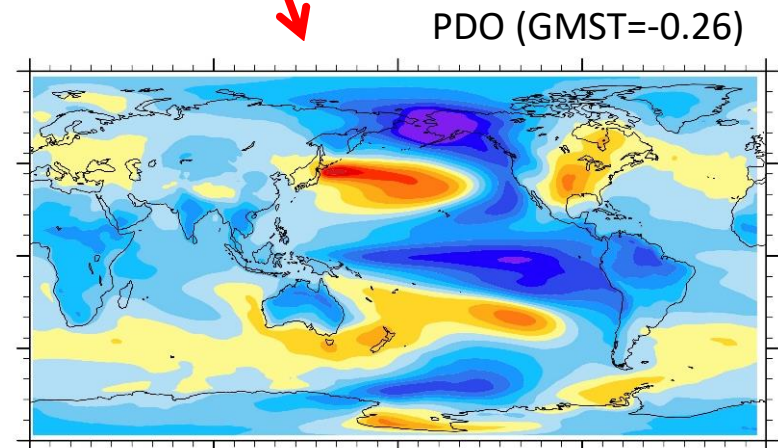


# Forced & internal changes in GMST are distinct in **pattern**, energetics, mechanism & predictability

$$T_{obs} = T_F + T_I$$



-0.6 -0.4 -0.2 0 0.2 0.4 0.6 0.8 1.0 1.25 1.5 1.75 2.5  
Trend ( $^{\circ}\text{C}$  over period)



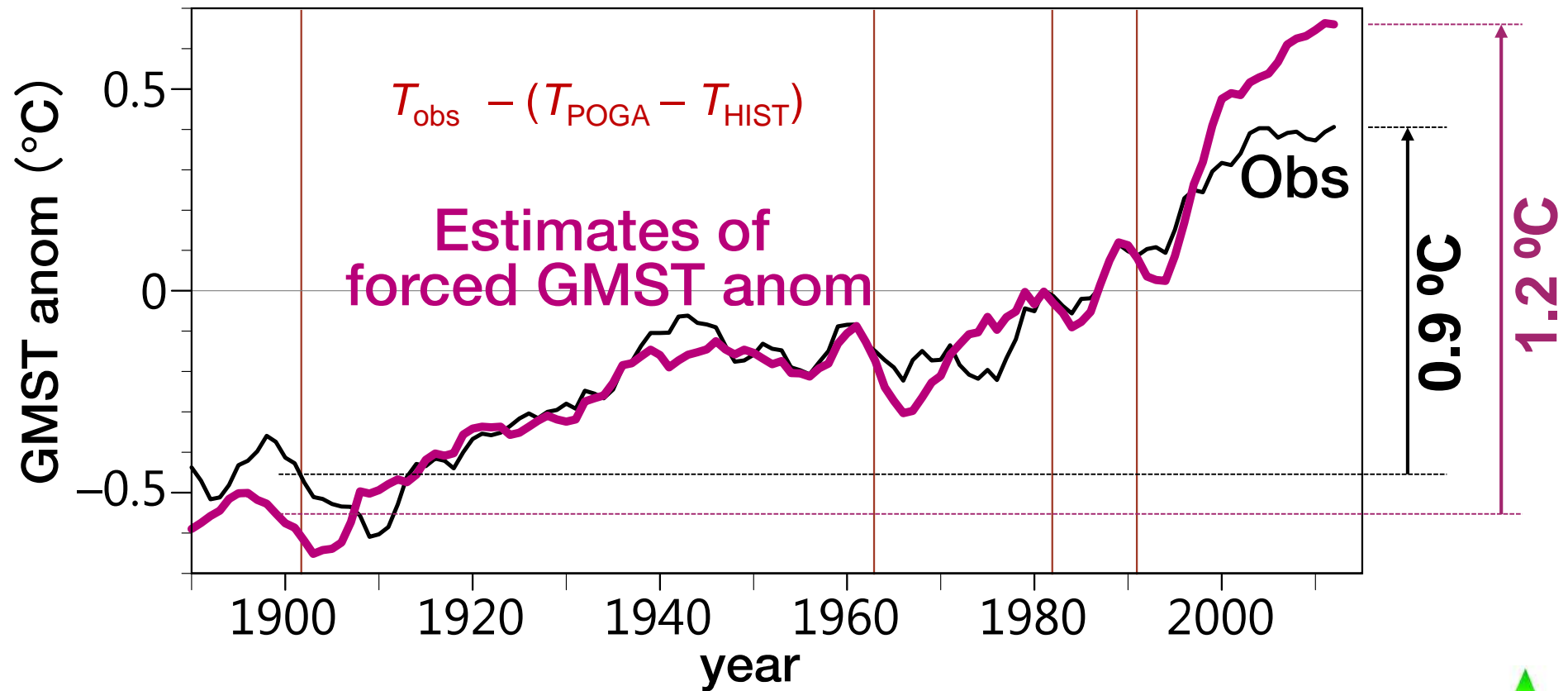
## Summary 3

- Tropical Pacific cooling is a major cause of the early 2000s hiatus.
- **PDO** dominates internal GMST variability, with **distinctive seasonal and spatial fingerprints.**
- Anthropogenic warming reached 1.2°C.

# Estimate anthropogenic warming

- Use **climate model HIST runs** (e.g., IPCC)
  - downside: uncertainties in **radiative forcing & climate feedback**
- Use **POGA pacemaker runs**
  - remove **internal variability** from the model (less affected by the above uncertainties):  $T_F = T_{obs} - T_I$

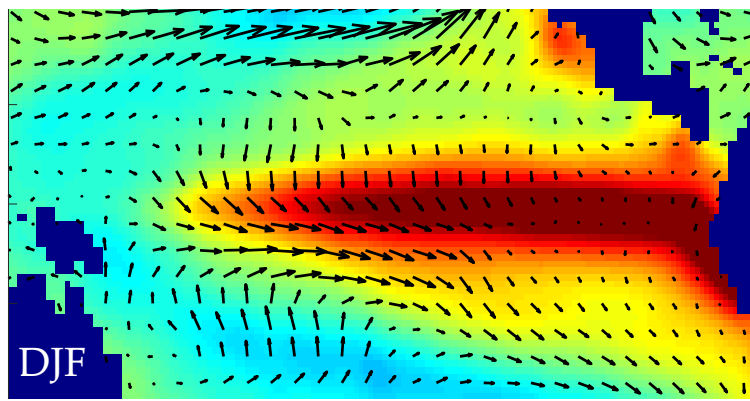
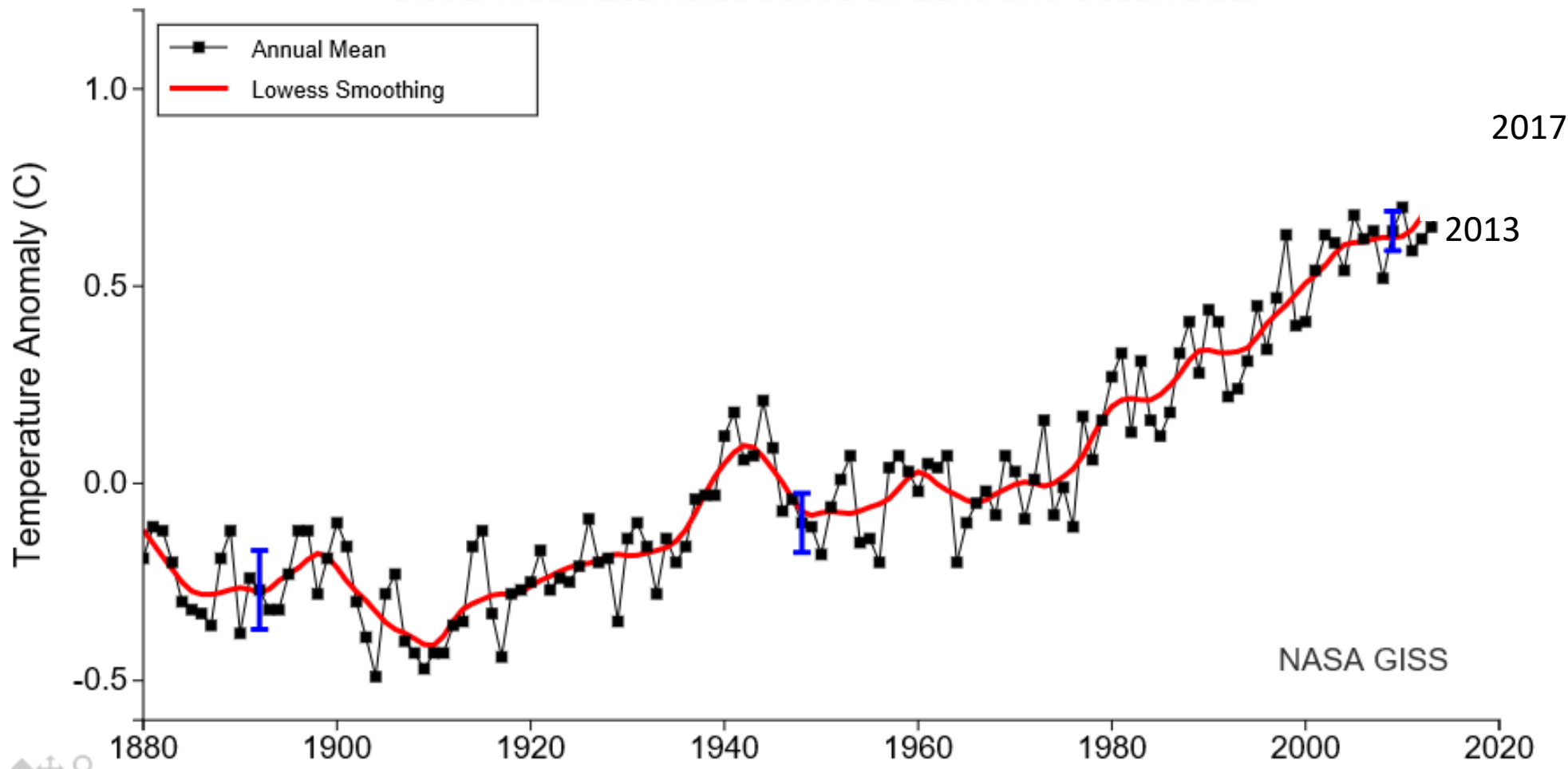
# Estimate anthropogenic warming from obs



- Paris agreement: *“holding GMST increase to well below 2°C above pre-industrial levels”*
- Anthropogenic warming already reached 1.2°C.



# Global Mean Estimates based on Land and Ocean Data



2015-16 El Niño

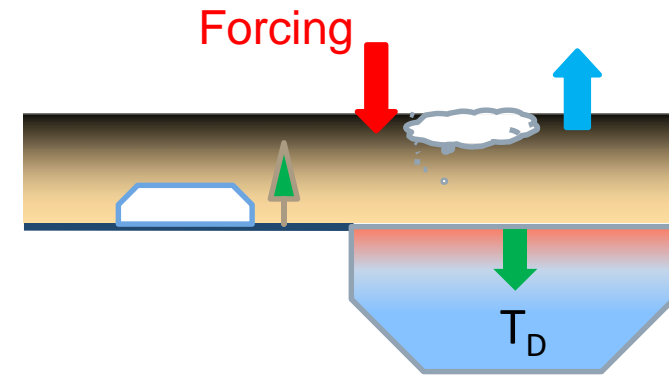
# Energy theory $H_t = NTO_A = F - \lambda T$

Ocean heat content change

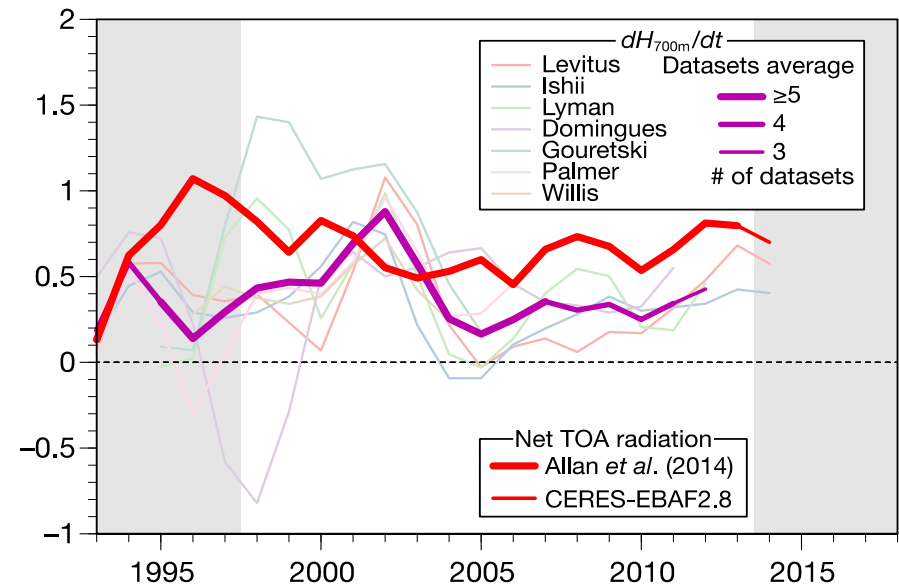
TOA radiative imbalance

Radiative forcing

Climate feedback



Prediction: during the hiatus (T=0), ocean heat uptake accelerates (TOA radiative imbalance increases).



TOA radiative imbalance (0.7 W/m<sup>2</sup>) from CERES satellite (2000-), calibrated against Argo data.